1. **Research Area**

Geophysical surveys were carried out in the area of the town of Działoszyce, located in the southern part of the Świętokrzyskie Voivodeship, within the Pińczów County. The town is situated at the confluence of the Jakubówka and Sancygniówka rivers, the latter being a tributary of the Nidzica River. This area is characterized by a long and well-documented history of settlement, dating back at least to the 13th century, as evidenced by the oldest written sources. As a result of the Third Partition of Poland in 1795, Działoszyce became part of the Habsburg Monarchy, remaining under Austrian rule for twelve years. Between 1807 and 1815, the town belonged to the Duchy of Warsaw, and from 1815 – following the resolutions of the Congress of Vienna – it was incorporated into the Russian Empire, where it remained for over a century. The proximity of the Austrian border and the neighbouring Free City of Kraków favored the development of settlement, including the intensification of Jewish settlement, particularly during the 19th century. During the Second Polish Republic, Działoszyce functioned as a local center for agricultural trade, dealing primarily in grain, poultry, and livestock. Several small industrial establishments also operated in the town. According to the 1921 census, Działoszyce had 6,755 inhabitants, approximately 80% of whom were Jewish. By 1939, the population had decreased to 5,872. During the German occupation, the Jewish community was subjected to extermination. Following the annexation of Upper Silesia into the Third Reich, large numbers of Jews and Poles relocated to the General Government, resulting in an increase in the population of the Działoszyce region. On September 2, 1942, Gestapo units from Miechów arrived in the town, ordering the gathering of all Jewish residents – both from Działoszyce and surrounding villages – in the town square. By the end of the day, approximately 10,000 individuals had assembled there. On September 3, orders were issued to transport the gathered population to a transit point near Miechów, and subsequently – primarily – to the Bełżec extermination camp. The majority of individuals were marched on foot to the railway station. About 1,500 people, including the elderly and those unable to walk, were loaded onto horse-drawn carts and trucks and transported to the Jewish cemetery, where they were executed. Their bodies were buried in three mass graves. In the following days, German soldiers searched homes abandoned by Jewish residents, confiscating and removing movable property. The deportations and extermination led to the near-total annihilation of the Jewish community in Działoszyce and a dramatic decrease in the town’s overall population – after the war, the number of inhabitants dropped to fewer than two thousand.

The geological structure of the near-surface layers in the Działoszyce area is dominated by Pleistocene aeolian deposits, mainly loess and silty covers. Their thickness ranges on average from several to over a dozen meters, and they exhibit considerable morphological variability at both regional and local scales. Complementing the geological setting are flat-bottomed valleys of the Jakubówka, Sancygniówka, and Nidzica rivers, filled with Holocene fluvial sands and silts forming floodplain terraces and valley floors. These deposits are bordered by glacial tills associated with the glacial complexes of southern Poland. Locally, the subsurface may also contain outcrops of Miocene gypsum or anhydrite, often in association with interbedded clays, marls, marly limestones, rock salt, and sulfur-bearing limestones – lithostratigraphic units classified as part of the Krzyżanowice Formation. In summary, the area covered by the geophysical survey is characterized by the predominance of low-resistivity loess and silty deposits, locally enriched with inclusions of glacial tills and alluvial sediments.

1. **Methodology**

This chapter presents the characteristics of two geophysical methods that have found widespread application in archaeological research – gradient magnetometry and electromagnetic conductivity (EM) surveying. Both techniques belong to the group of non-invasive geophysical prospection methods and enable the identification of subsurface structures with varying physical properties, without the need for intrusive excavation. Their complementary nature allows for the effective detection of both anomalies related to anthropogenic modifications of the soil environment and natural lithological variations. The following sections of the chapter discuss the physical principles underlying each method, their range of applications, procedures for data acquisition, and the specific aspects of data interpretation in the context of archaeological investigations.

* 1. **Magnetometry method**

Magnetometry is a non-invasive, surface-based geophysical method that measures the intensity of the Earth's magnetic field. Variations in this field make it possible to detect different magnetic properties of the subsurface, which has led to applications across various disciplines. In geology, it is used to identify magmatic structures, while in engineering it is applied in the investigation of pipelines and the detection of unexploded ordnance. In archaeology, magnetometry enables the mapping of buried structures, hearths, and even graves.

Objects located within a magnetic field are subject to the phenomenon of induction, meaning they become magnetized. The degree of magnetization depends on the intensity vector T and the magnetic susceptibility of the material, which allows for the classification of substances into different categories. Diamagnetic, paramagnetic, and antiferromagnetic materials all exhibit different levels of magnetization. However, ferromagnetic materials interact most strongly with the magnetic field, causing measurable disturbances in its intensity. These anomalies are precisely what magnetometric prospection seeks to detect. Using this method, it is possible to infer the presence of various ferromagnetic structures and objects within the Earth’s crust. The discovery of such anomalies may indicate the presence of ore deposits, crystalline mineral bodies, magmatic veins, or archaeological and historical relics. Magnetometry is also useful in detecting components of underground infrastructure and utility networks.

In this study, a specific variant of the magnetometric method was applied—measurement of the vertical magnetic gradient. This technique involves the simultaneous measurement of the magnetic field using two sensors placed at different vertical positions. The greater the distance between the sensors, the deeper the investigation range of the method, allowing for detection at greater depths. The primary application area of this technique is archaeological research. It is a dedicated tool that enables the efficient identification of buried structures and hearths without the need for simultaneous reference measurements. As a result, the survey process is simplified, and the obtained data are both accurate and precise.

* 1. **Slingram (FDEM)**

The electromagnetic conductivity method (*FDEM or slingram*) is one of the electromagnetic techniques used in geophysical surveys. The functioning of a conductivity meter relies fundamentally on the generation, transmission, and reception of electric current. The instrument utilizes two coils, spaced at a distance s from each other. The first of these, the transmitter coil, generates a low-frequency alternating current (typically in the range of several to several tens of kilohertz). When current flows through the transmitter coil, it produces a time-varying magnetic field, referred to as the primary magnetic field (Hp). This primary field plays a key role, as it penetrates the subsurface medium. As a result, eddy currents are induced in the ground, giving rise to a secondary magnetic field.

The receiver coil, placed at a fixed distance from the transmitter, registers this secondary magnetic field generated in the subsurface. This process operates according to the principle of magnetic induction, which induces a voltage (and thereby a current) in the receiver. There is a direct correlation between the electrical conductivity of the subsurface (as affected by magnetic induction) and the difference in magnetic field strength between the primary field generated by the transmitter coil (**Hp**) and the secondary field measured by the receiver coil (**Hs**).

In the present study, a conductivity meter equipped with six transmitter–receiver coil pairs was used, allowing for simultaneous measurements of ground conductivity at six different investigation depths. Two key parameters are recorded during the survey: the apparent conductivity, which reflects variations in the electrical properties of the ground, and the relative amplitude of the vertical component of the magnetic field. This second parameter - known as the *in-phase* response - enables the detection of subsurface metallic objects.

1. **Processing**

Both methods require initial processing of the recorded measurement data, followed by proper interpretation in both geophysical and archaeological contexts. A detailed description of the individual stages of data processing will be presented separately for each of the discussed methods. Measurements using both techniques were carried out at five survey sites located in close proximity to one another. However, each of these areas represented a distinct interpretative challenge.

Three of the survey polygons were situated in the immediate vicinity of the Jewish Cemetery memorial in Działoszyce, while two additional survey areas were located at some distance, on opposite sides of a forest road. The survey terrain was characterized by a high degree of complexity in terms of its suitability for the applied geophysical methods. For the first three polygons, a major source of interference was the nearby power line. Additionally, metal fencing surrounding the memorial - located in close proximity to polygons 2 and 3 - had a significant impact on data quality. These areas exhibited the highest level of signal disturbance for both measurement methods.

The fourth polygon was located within a loess ravine on one side of the road, whereas the fifth was situated in a small depression on the opposite side. At these two sites, the main sources of interference were scattered metal debris such as wires, cans, old utensils, and bottle caps. Where possible, the area was cleared of surface debris prior to the commencement of measurements.

* 1. **Magnetometry**

The survey was conducted using a quasi-regular measurement grid along profiles spaced approximately 50 cm apart, taking into account areas where measurements could not be performed due to dense vegetation or the presence of trees. Data processing was carried out using standard procedures. Distorted or saturated readings, as well as isolated point anomalies (so-called "spikes"), were excluded from further processing. Such point anomalies were defined as values exceeding 1000 nT/m relative to both neighboring measurement points.

Subsequently, the cleaned dataset was interpolated using kriging, with simultaneous exclusion of areas lacking data coverage. In the visualizations, zones of positive vertical magnetic gradient are marked in orange, while negative values are shown in blue.

The vertical magnetic gradient is calculated as the difference between the reading from the lower sensor and that from the upper sensor. In this study, the vertical spacing between the sensors was set to 1 meter, as the focus was on detecting near-surface features. A positive magnetic response is expected directly above zones containing ferromagnetic materials, while a negative response typically appears at some distance from such sources. Due to the proximity of metallic elements, such as the aforementioned fence, extreme values observed on the gradient maps should be attributed primarily to these surface or near-surface interferences.

* 1. **Slingram (FDEM)**

The survey was conducted using a quasi-regular measurement grid, including areas where data acquisition was not possible due to dense undergrowth or the presence of trees. Data processing followed standard procedures. Distorted or saturated measurements, as well as isolated point anomalies (commonly referred to as "spikes"), were excluded from further analysis.

A one-dimensional (1D) inversion of the resistivity data was carried out, whereby for each field curve a model curve was calculated to minimize the misfit. The maximum acceptable error was set at 5%. Each measurement point was treated as an individual sounding, and the 1D interpretation was performed separately for every location.

The resulting data were interpolated using kriging, with areas lacking data coverage excluded from the final interpolation. In the visualizations, zones of high resistivity are shown in red, while zones of low resistivity are marked in blue.

1. **Results**

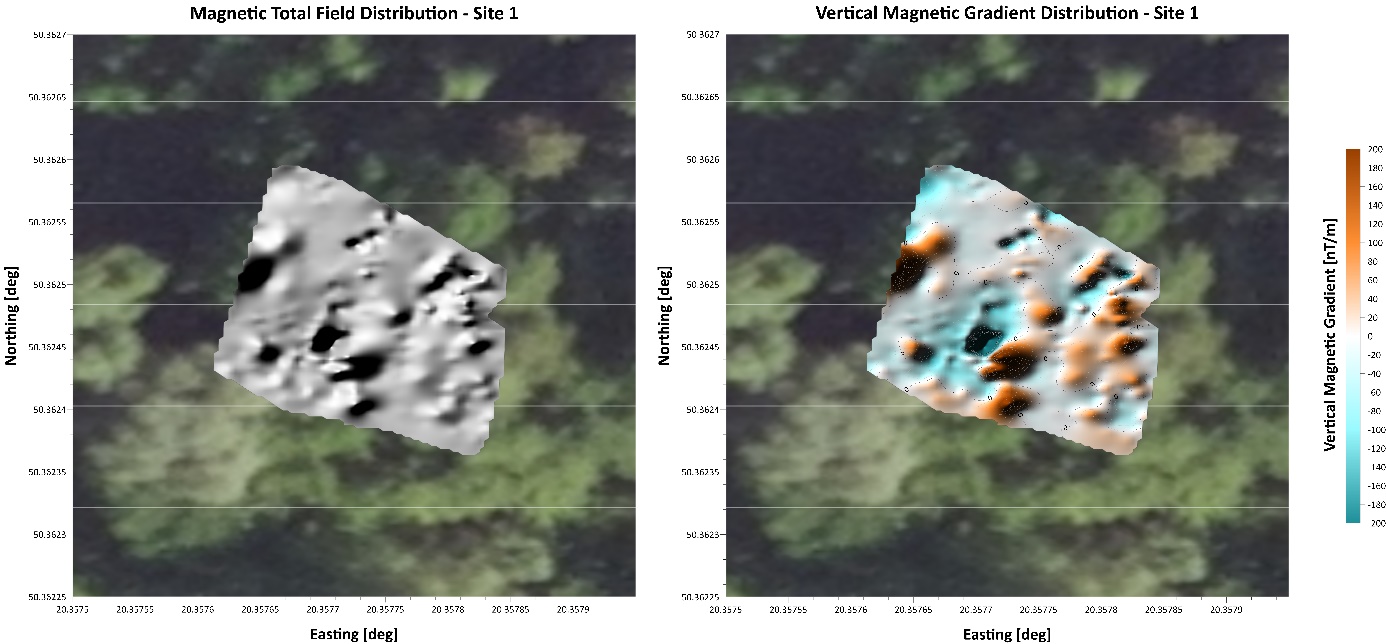
This chapter presents the results of geophysical measurements conducted across five designated survey polygons. Each area is analyzed individually, following a predetermined sequence. The presentation begins with maps obtained from magnetometric surveys. For each area, a pair of maps is shown: on the left, the distribution of the total magnetic field visualized on a shaded relief map, and on the right, the map of the vertical magnetic gradient. To eliminate the influence of global field variations and to focus solely on residual anomalies, the total magnetic field data were normalized in advance. The normalization process was based on reference data obtained from the Central Geophysical Observatory in Belsk, made available through the *Intermagnet* network.

Subsequently, for each area, an analogous set of maps is provided, indicating potential structures that may correspond to burial sites. These maps aim to visualize possible interpretations of the geophysical anomalies in the context of the targeted archaeological features.

Following the same structure, results obtained using the electromagnetic conductivity method are presented. For each analysed area, a series of six maps illustrates the distribution of electrical conductivity (the inverse of resistivity) at selected investigation depths: 0.25 m, 0.5 m, 0.8 m, 1.1 m, 1.6 m, and 2.3 m. Potential locations of subsurface features interpreted as burial-related structures are then marked on each of these maps. This presentation format allows for a comparative analysis of results obtained using two independent methods and facilitates the assessment of spatial consistency in the interpretation.

* 1. **Area 1**

Figures 1.1 and 1.2 present the results of magnetometric measurements conducted in the area designated as Polygon 1. On the map of the total magnetic field, a clearly defined, quasi-rectangular anomaly is visible. In its central part, this feature coincides with a characteristic transition zone between a positive and a negative maximum on the map of the vertical magnetic gradient. This configuration may indicate the presence of a disturbed subsurface structure - most likely a linear feature such as a trench - initially interpreted as a potential mass grave. Notably, the orientation of the anomaly is perpendicular to the direction of the survey profiles, which reduces the likelihood of the so-called "striping effect" caused by acquisition directionality. In other parts of the polygon, numerous point anomalies are observed, though their origins remain ambiguous. These suggest a high level of signal disturbance and noise within the surveyed area.



*Fig.1.1. Distribution of the total magnetic field (left) and the vertical magnetic gradient (right) within the boundaries of survey area no. 1.*

Obraz zawierający tekst, zrzut ekranu

Zawartość wygenerowana przez sztuczną inteligencję może być niepoprawna.

*Fig.1.2. Distribution of the total magnetic field (left) and the vertical magnetic gradient (right) within the boundaries of survey area no. 1,* *with the magnetic anomaly interpreted from magnetometric data marked in red and the anomaly interpreted from conductivity data marked in green.*

Figures 1.3 and 1.4 present the distribution of electrical resistivity for Polygon 1, obtained using the electromagnetic conductivity method. In the area corresponding to the previously identified magnetic anomaly, a dominant resistivity structure is clearly visible. This feature is characterized by low resistivity values in the near-surface layers—down to approximately 1 meter in depth -followed by a transition into a zone of elevated resistivity. This pattern stands in clear contrast to the surrounding substrate, where resistivity variation with depth is significantly less pronounced.

The interpretation of this anomaly may suggest the presence of an anthropogenic feature -potentially a mass grave - in which human remains are located below the 1-meter depth, with the overlying space filled by highly conductive material, such as locally occurring clays or fluvial silts. This type of resistivity contrast may result from both the physical properties of the backfill material and structural disturbances within the natural lithological setting.

***Obraz zawierający zrzut ekranu, tekst

Zawartość wygenerowana przez sztuczną inteligencję może być niepoprawna.***

*Fig.1.3. Distribution of subsurface electrical resistivity at depths of 0.25, 0.5, 0.8, 1.1, 1.6, and 2.3 m within the boundaries of survey area no. 1.*

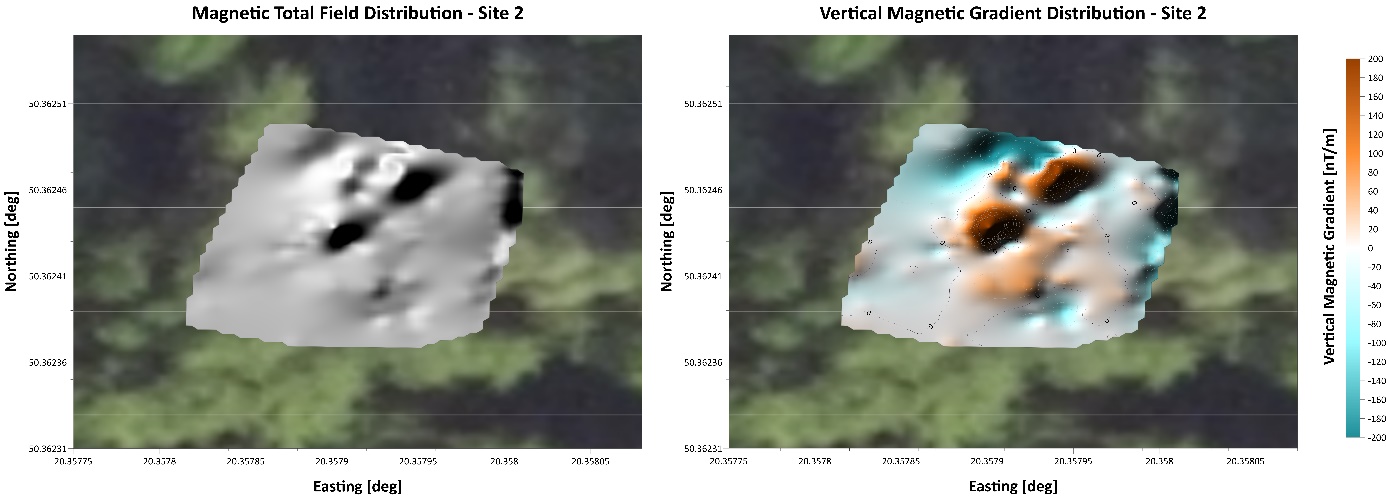
*Obraz zawierający zrzut ekranu, tekst

Zawartość wygenerowana przez sztuczną inteligencję może być niepoprawna. Fig.1.4.* *Distribution of subsurface electrical resistivity at depths of 0.25, 0.5, 0.8, 1.1, 1.6, and 2.3 m within the boundaries of survey area no. 1, with the magnetic anomaly interpreted from magnetometric data marked in red and the anomaly interpreted from conductivity data marked in green.*

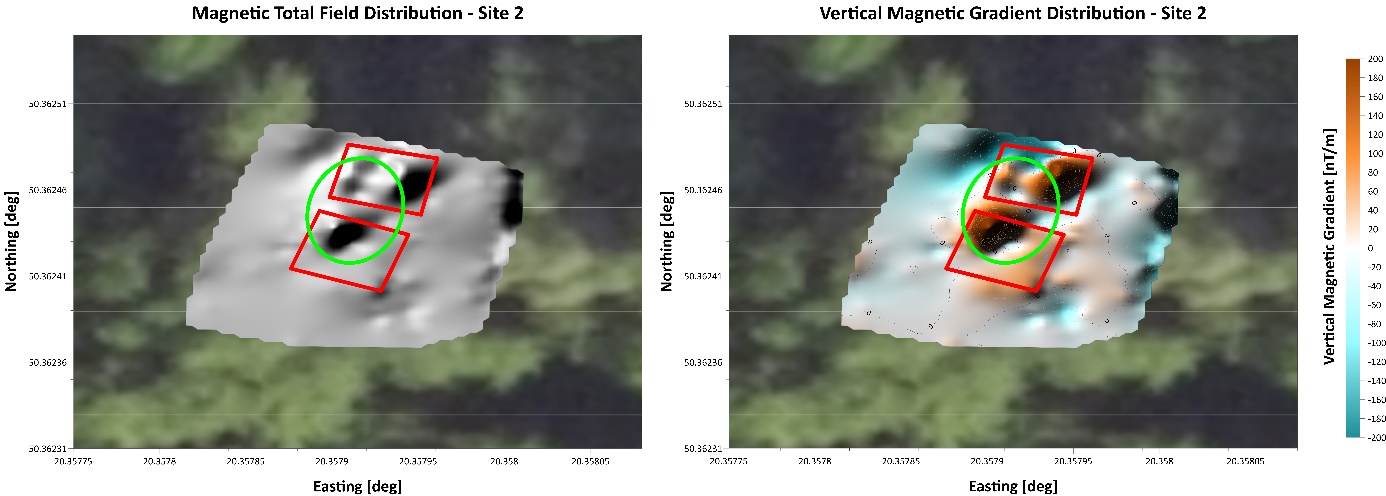
* 1. **Area 2**

Figures 2.1 and 2.2 present the results of magnetometric measurements conducted within the area designated as Polygon 2. In the northern part of the surveyed area, an anomalous structure is visible, consisting of two closely spaced maxima of the vertical magnetic gradient, separated by a local minimum. This pattern is also reflected on the map of the total magnetic field; however, the structure as a whole appears less distinct and more difficult to interpret compared to the anomaly observed in Polygon 1.

Additionally, a strong magnetic effect is visible in the eastern part of the polygon, which can be clearly attributed to the presence of a metal fence surrounding the nearby monument, located in close proximity to the measurement area. Compared to Polygon 1, the number of point anomalies interfering with interpretation is lower, resulting in a slightly higher overall clarity of the data. It is worth noting that Polygons 1 and 2 are directly adjacent to each other, allowing for their spatial correlation in the subsequent analysis.

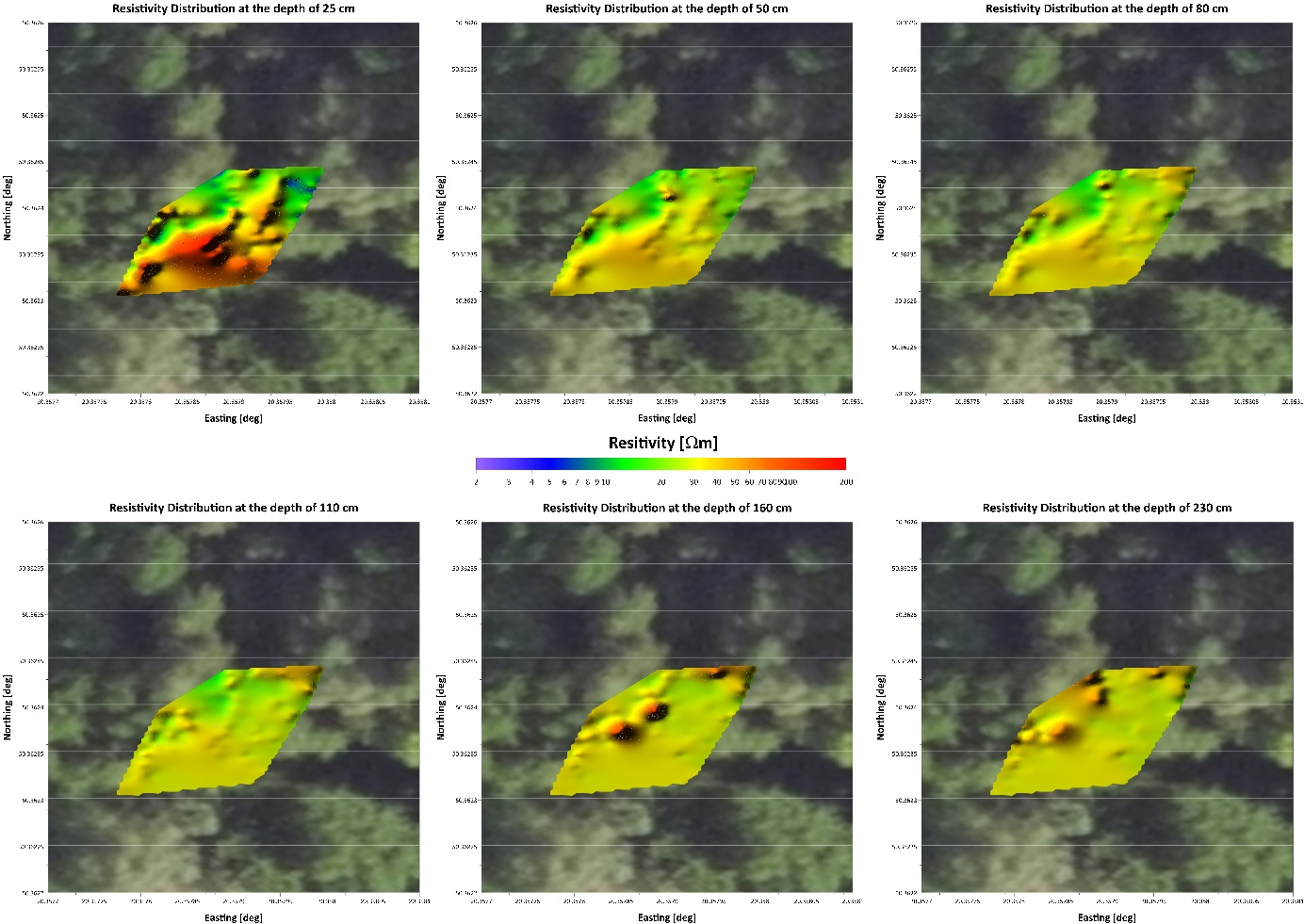


*Fig.2.1. Distribution of the total magnetic field (left) and the vertical magnetic gradient (right) within the boundaries of survey area no. 2.*

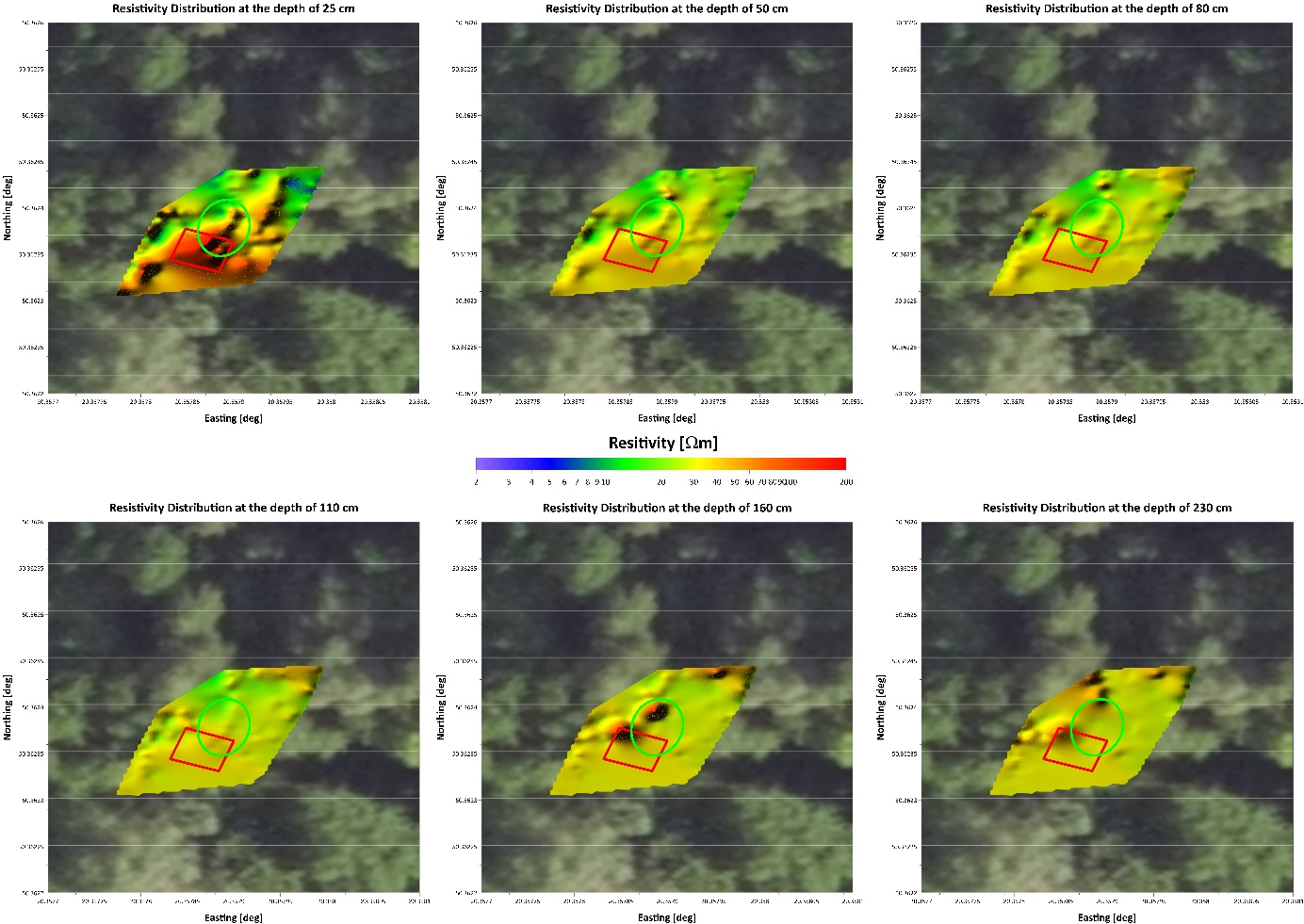


*Fig.2.2. Distribution of the total magnetic field (left) and the vertical magnetic gradient (right) within the boundaries of survey area no. 2, with the magnetic anomaly interpreted from magnetometric data marked in red and the anomaly interpreted from conductivity data marked in green.*

A different picture compared to the magnetic data is presented by the distribution of electrical resistivity, as shown in Figures 2.3 and 2.4. In this dataset, a clear linear anomaly is visible in the eastern part of the polygon, which can be confidently associated with the course of the metal fence—previously identified as a source of interference in the magnetometric measurements.   
A pronounced division of the surveyed area is especially evident in the shallower depth intervals: the southern part of the polygon is characterized by higher resistivity values, while the northern part shows lower values. This contrast gradually diminishes with increasing depth.

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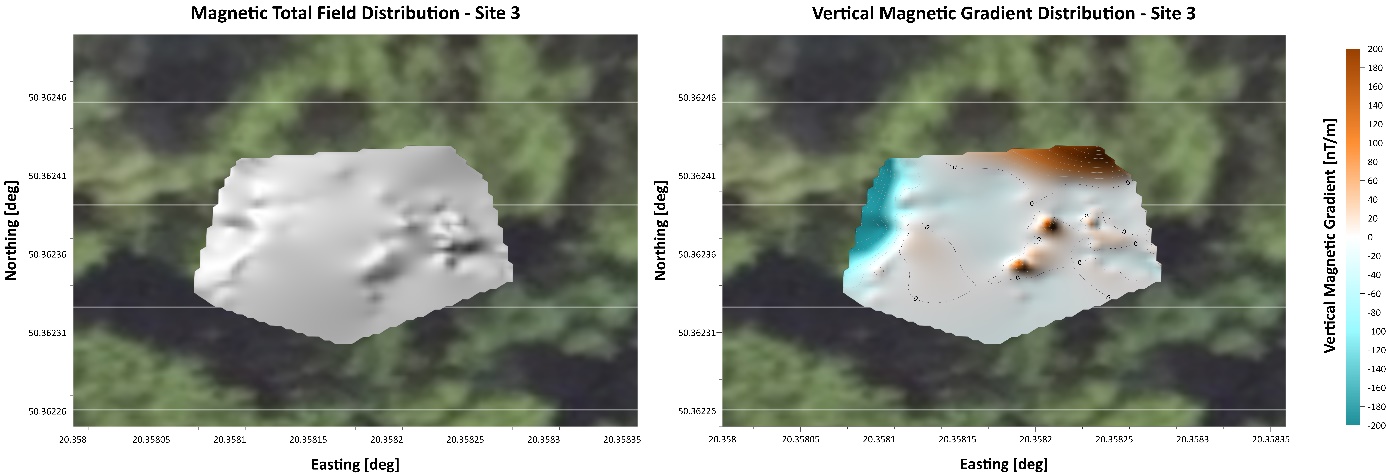
*Fig.2.3. Distribution of subsurface electrical resistivity at depths of 0.25, 0.5, 0.8, 1.1, 1.6, and 2.3 m within the boundaries of survey area no. 2.*

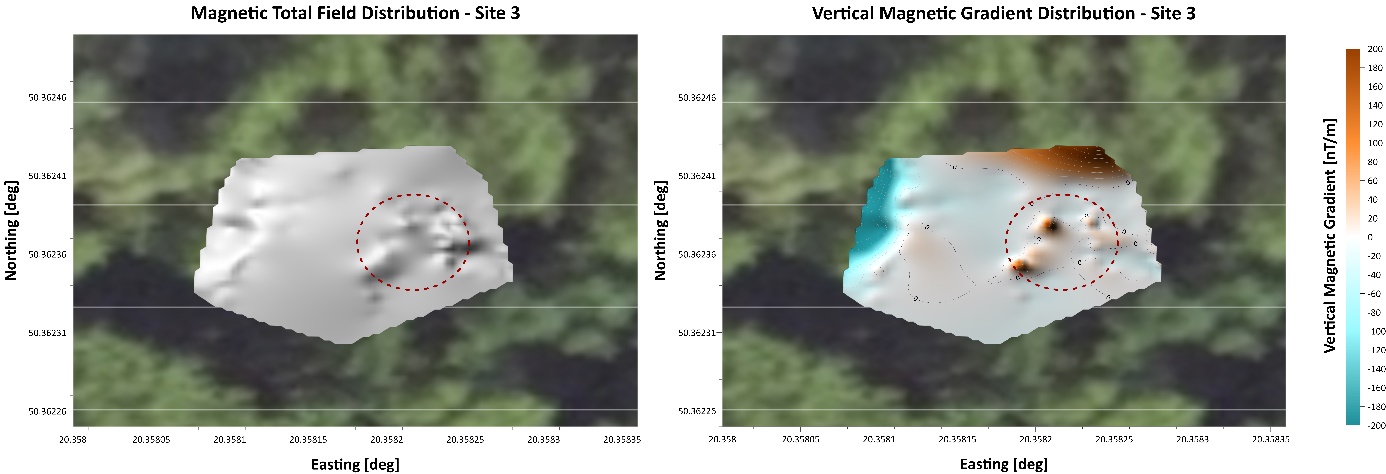
* Fig.2.4. Distribution of subsurface electrical resistivity at depths of 0.25, 0.5, 0.8, 1.1, 1.6, and 2.3 m within the boundaries of survey area no. 2, with the magnetic anomaly interpreted from magnetometric data marked in red and the anomaly interpreted from conductivity data marked in green.*

* 1. **Area 3**

The results of magnetometric measurements for Polygon 3, presented in Figures 3.1 and 3.2, appear to be relatively straightforward in terms of interpretation. The dominant anomaly is located in the western part of the area, where a strong magnetic response is observed - clearly associated with the presence of a metal fence surrounding the monument. In the central part of the polygon, a cluster of positive point anomalies is visible on the vertical magnetic gradient map, partially reflected in the irregular pattern of the total magnetic field distribution.

This structure has been marked with a maroon dashed circle, indicating an area of ambiguous interpretative character - there is a lack of clear evidence for determining its origin, although its potentially anthropogenic nature cannot be excluded.

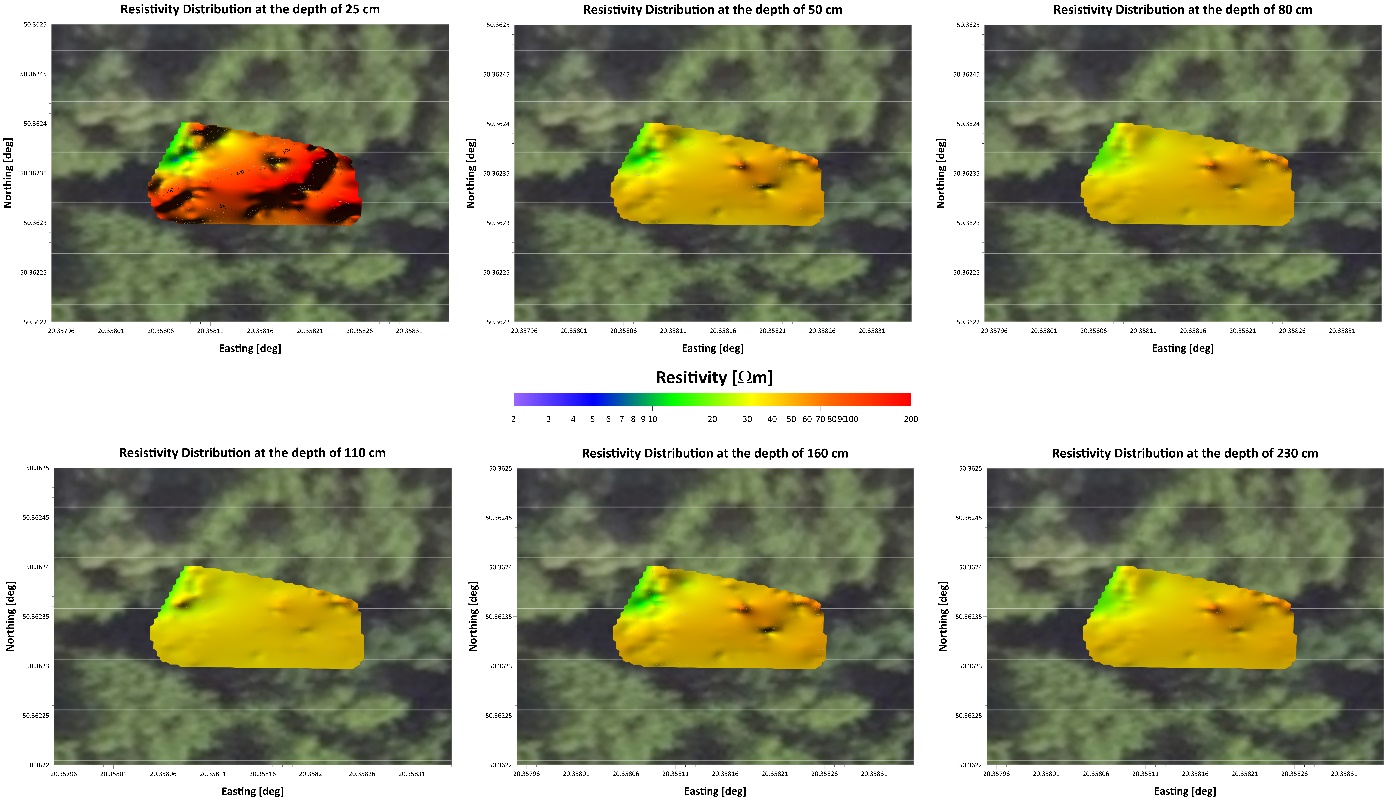
*Fig.3.1. Distribution of the total magnetic field (left) and the vertical magnetic gradient (right) within the boundaries of survey area no. 3.*



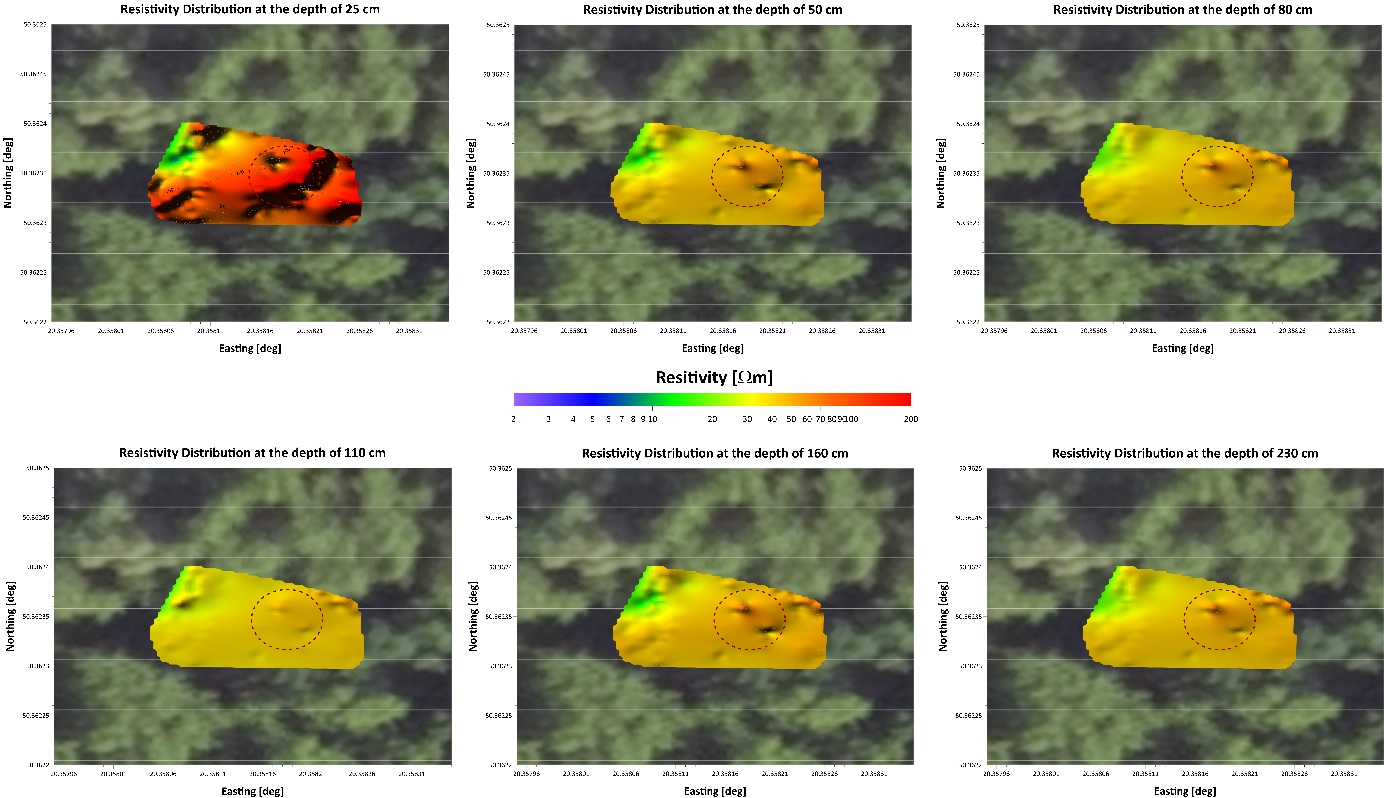
*Fig.3.2. Distribution of the total magnetic field (left) and the vertical magnetic gradient (right) within the boundaries of survey area no. 3, with the magnetic anomalies marked in dashed red lines.*

Figures 3.3 and 3.4 show the distribution of electrical resistivity for Polygon 3. Similar to the magnetometric data, the dominant interpretative feature is a strong anomaly located in the western part of the area, resulting from the influence of the metal fence surrounding the monument. Notably, the near-surface resistivity values are significantly higher than those observed in the previously analysed polygons, despite their close spatial proximity. This may indicate local variations in lithological properties or moisture conditions.

In the eastern part of the polygon, a cluster of point anomalies -local resistivity maxima - can be observed. The origin of these anomalies remains unclear. While their spatial distribution suggests a certain degree of coherence, there is no clear correlation with the magnetometric data or the archaeological context, making their interpretation ambiguous.

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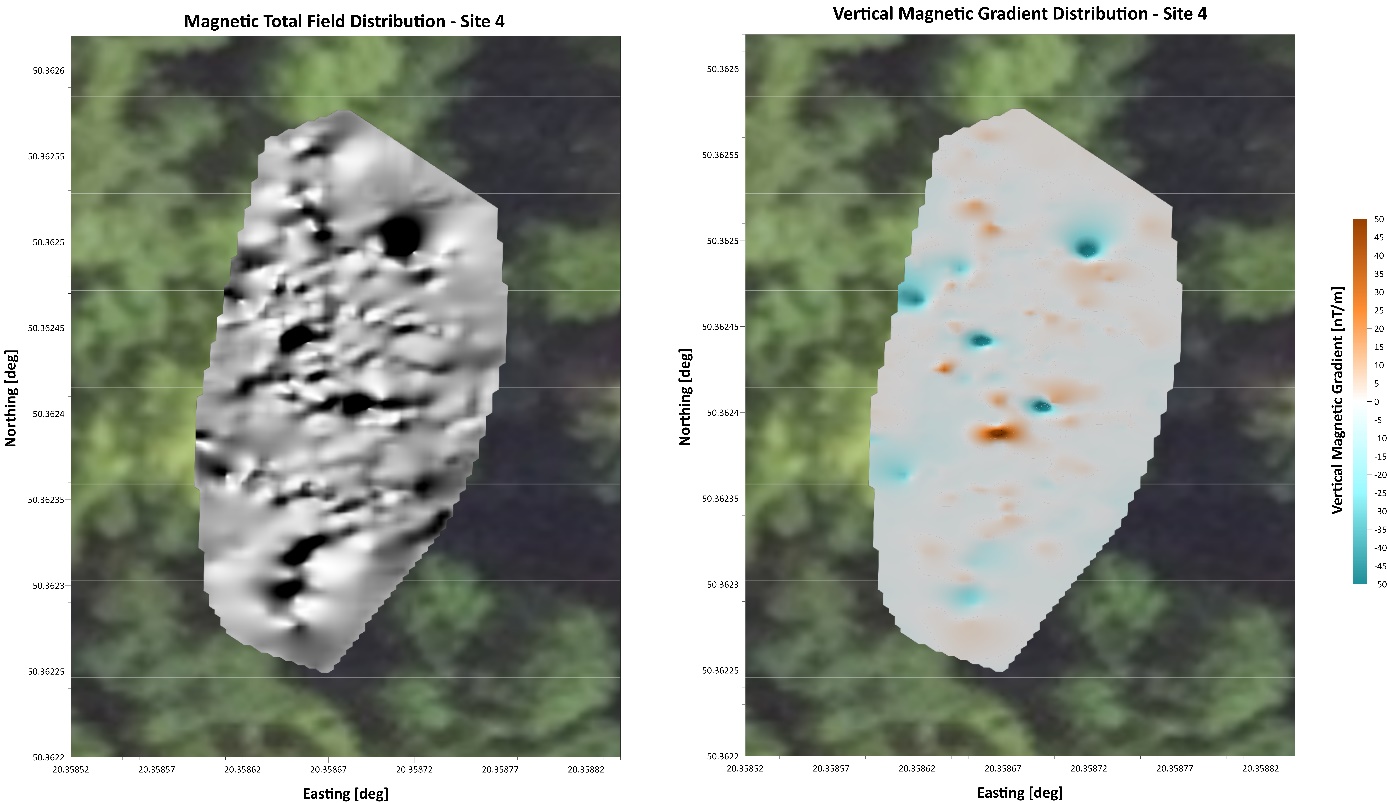
*Fig.3.3. Distribution of subsurface electrical resistivity at depths of 0.25, 0.5, 0.8, 1.1, 1.6, and 2.3 m within the boundaries of survey area no. 3.*

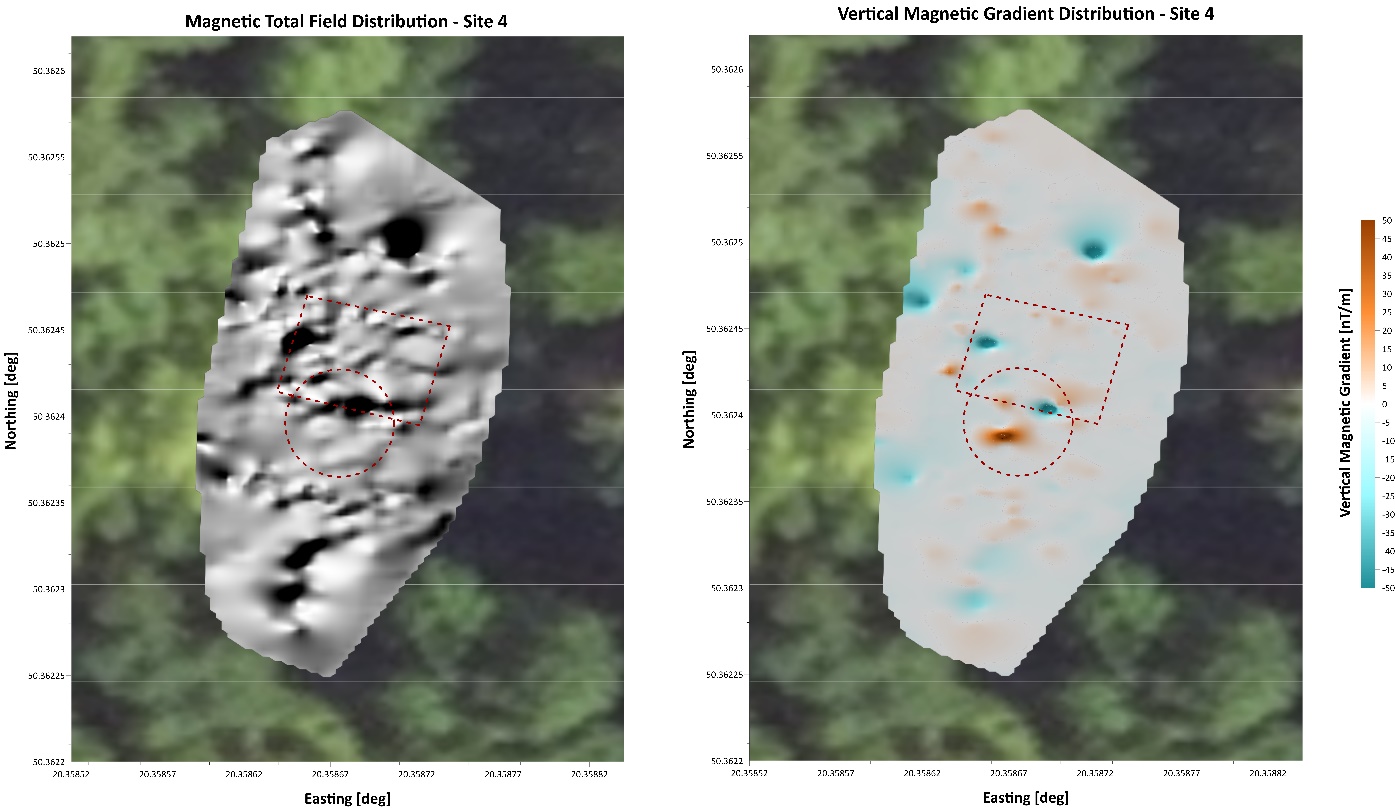
* Fig.3.4.* *Distribution of subsurface electrical resistivity at depths of 0.25, 0.5, 0.8, 1.1, 1.6, and 2.3 m within the boundaries of survey area no. 3, with the magnetic anomalies marked in dashed red lines.*

* 1. **Area 4**

The results of magnetometric measurements for Polygon 4 are presented in Figures 4.1 and 4.2. One of the key features distinguishing this area from the previously analyzed polygons is the noticeably weaker magnetic response. The values of the vertical magnetic gradient range from -50 to +50 nT/m, which contrasts with the broader scale of -200 to +200 nT/m used in the data visualizations for Polygons 1 - 3. This lower signal intensity results in generally weaker contrast between magnetic anomalies in this area.

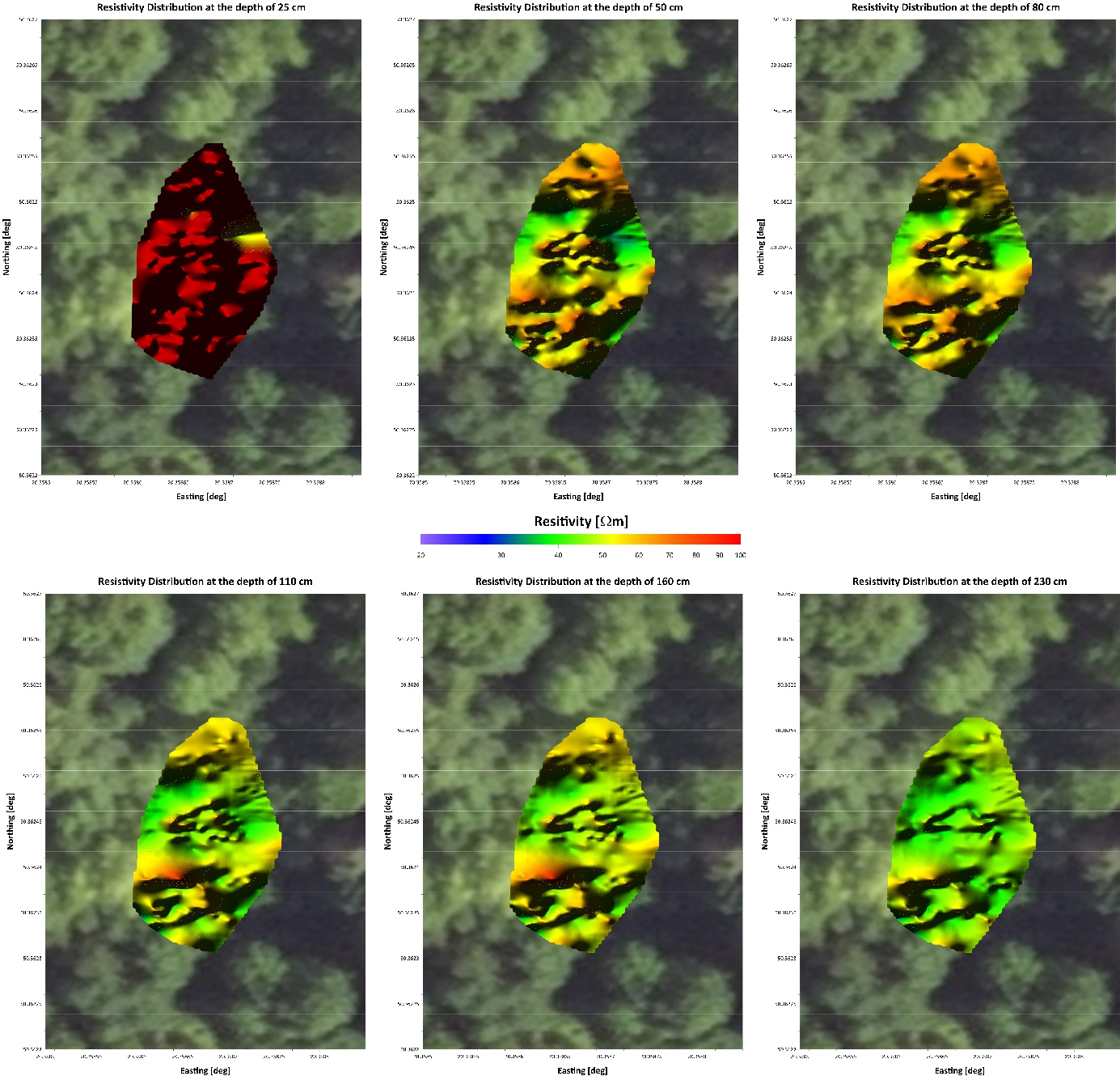
In the central part of the polygon, a small cluster of local maxima and minima in the magnetic gradient was recorded, which partially corresponds with linear structures visible on the total magnetic field map. However, this correlation is not clear-cut, and the interpretation of these features is far less distinct than in the case of, for example, Polygon 1. A potential area indicating the presence of an anthropogenic structure has been marked with a maroon dashed rectangle and circle, identifying a zone that requires further verification.

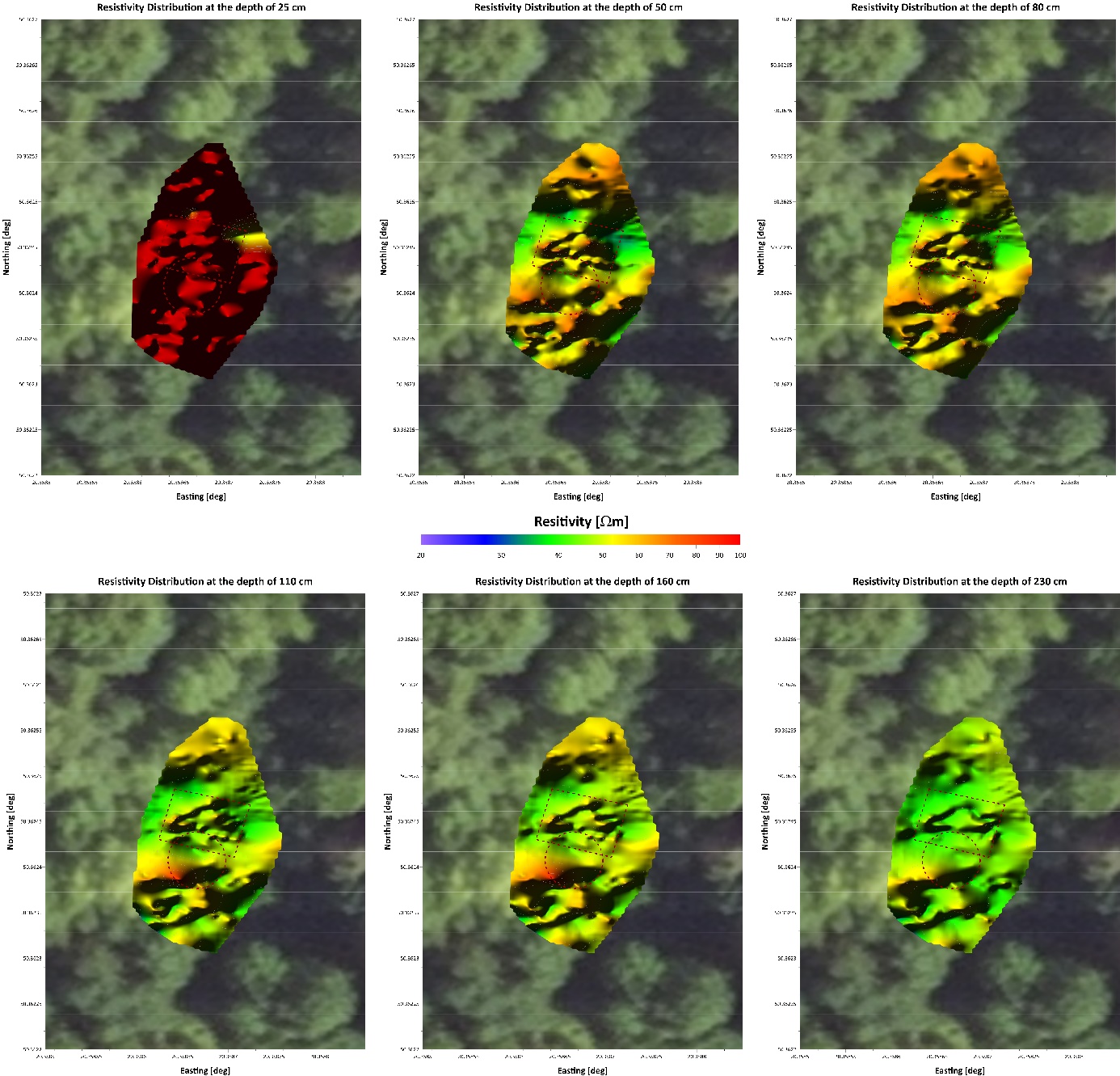
*Fig.4.1. Distribution of the total magnetic field (left) and the vertical magnetic gradient (right) within the boundaries of survey area no. 4.*

*Fig.4.2. Distribution of the total magnetic field (left) and the vertical magnetic gradient (right) within the boundaries of survey area no. 4, with the magnetic anomalies marked in dashed red lines.*

In the case of Polygon 4, the most promising results were obtained using the electromagnetic conductivity method. These results are presented in Figures 4.3 and 4.4. Excluding the near-surface layer - which exhibited extremely high resistivity values, most likely due to the drying of loess deposits - deeper layers revealed a distinct linear pattern of resistivity variation, oriented transversely to the axis of the ravine. The northern and southern parts of the area are dominated by high-resistivity structures, while in the central part, a zone of lower electrical resistivity emerges, morphologically resembling a channel or trench.

Within this central zone, at depths of up to approximately 1 meter, a structure with elevated resistivity is observed - comparable to that of the surrounding layers. Below this level, resistivity decreases again, approaching the values of the geological background. This phenomenon may indicate the presence of a filled void (e.g., a cavity or trench) that was backfilled with material originating from the surrounding area - a scenario consistent with anthropogenic infilling. Notably, the spatial position of this structure corresponds with the previously described magnetic anomaly, reinforcing the hypothesis of a potential archaeological feature. At the same time, due to the specific topography and genesis of the ravine, the influence of natural denudation processes on the observed resistivity pattern cannot be ruled out.

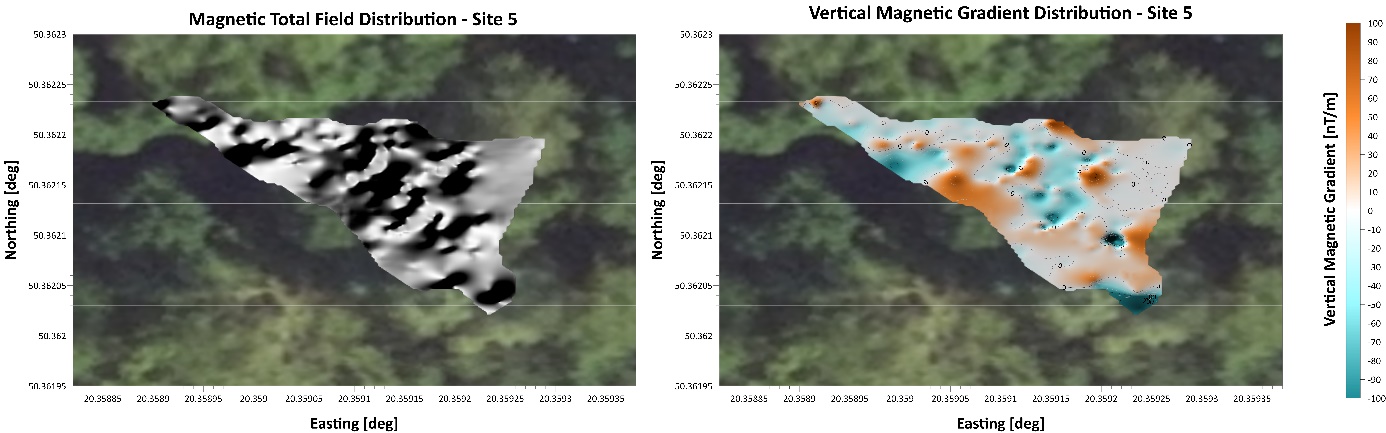
*******Fig.4.3. Distribution of subsurface electrical resistivity at depths of 0.25, 0.5, 0.8, 1.1, 1.6, and 2.3 m within the boundaries of survey area no. 4.*

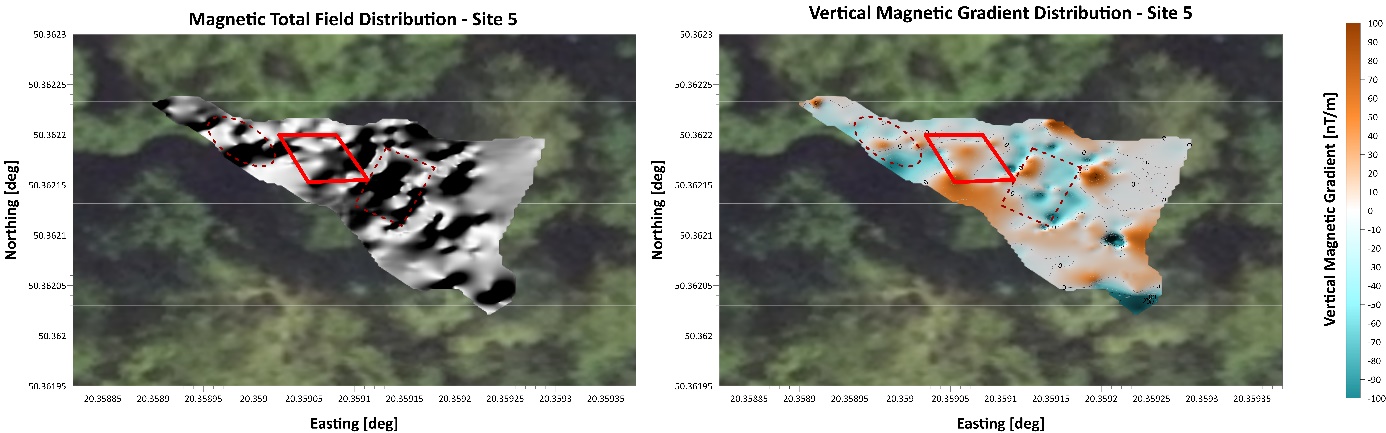
* Fig.4.4. Distribution of subsurface electrical resistivity at depths of 0.25, 0.5, 0.8, 1.1, 1.6, and 2.3 m within the boundaries of survey area no. 4, with the magnetic anomalies marked in dashed red lines.*

* 1. **Area 5**

Polygon 5 was located near a forest road, specifically at the junction of two forest paths, which gave it an unusual triangular shape. The results of the magnetometric measurements are presented in Figures 5.1 and 5.2. Compared to Polygon 4, the recorded anomalies are stronger; however, it should be noted that this area was heavily contaminated with surface debris - numerous scattered metal objects were present, including fragments of wire, pots, bottles, and other items capable of generating magnetic interference.

Along the northern edge of the polygon, adjacent to the road, a cluster of point-like maxima and minima in the magnetic gradient is visible. However, interpreting the origin of these anomalies is difficult. The area is also characterized by an irregular and ambiguous pattern in the distribution of the total magnetic field. Due to the presence of numerous disturbances and the overall uncertainty of the signal, the results from this polygon should be treated with caution - as supplementary data requiring validation through additional methods.

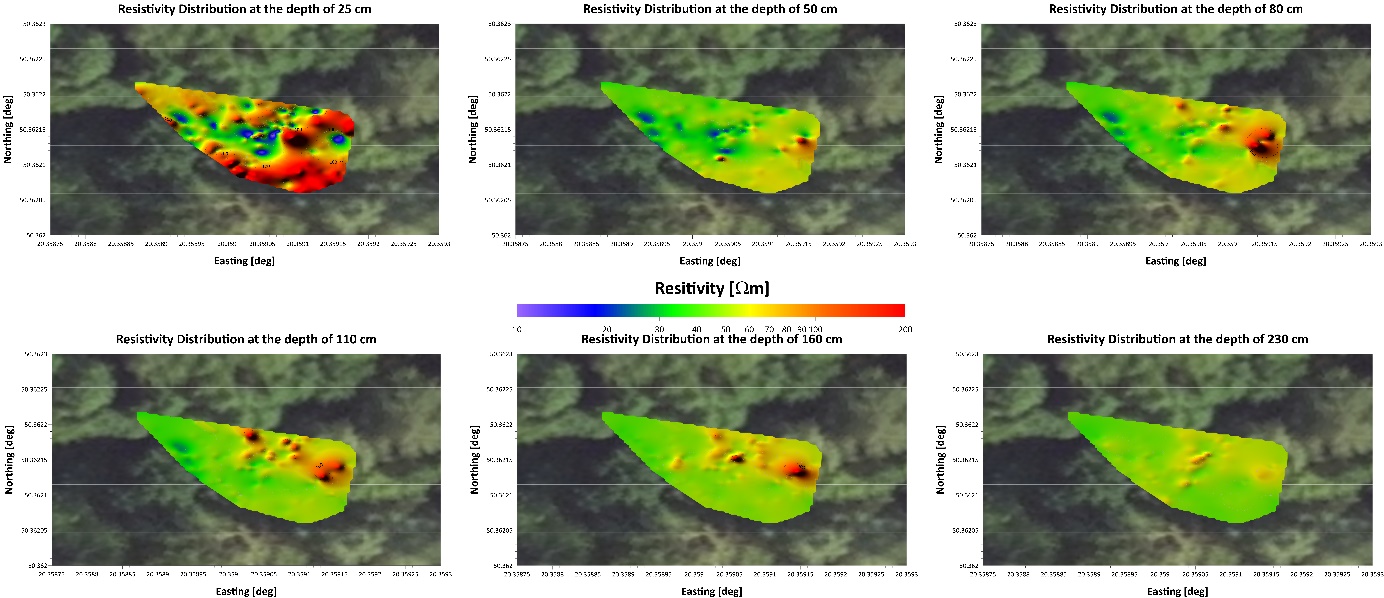
*Fig.5.1. Distribution of the total magnetic field (left) and the vertical magnetic gradient (right) within the boundaries of survey area no. 5.*



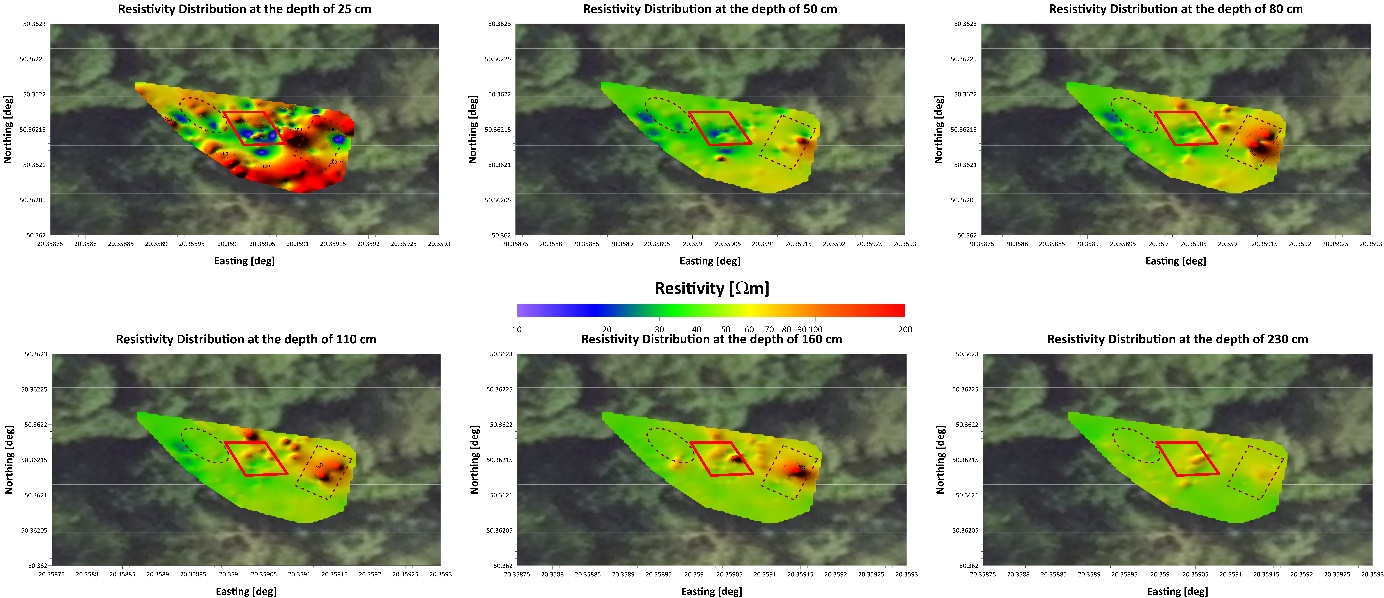
*Fig.5.2.* *Distribution of the total magnetic field (left) and the vertical magnetic gradient (right) within the boundaries of survey area no. 5, with the magnetic anomalies marked in dashed red lines.*

The results of conductivity measurements for Polygon 5 are presented in Figures 5.3 and 5.4. The most prominent feature in the resistivity data is a high-resistivity zone located in the northeastern part of the polygon, interpreted as an artificially formed earthen mound situated adjacent to the road. This structure is particularly well defined at depths ranging from approximately 0.8 to 1.6 meters below the ground surface, but it is not clearly visible in the deepest measurement range.

In other parts of the area, local point anomalies with low resistivity values were recorded at depths of up to about 0.5 meters. Some of these anomalies show partial correlation with the magnetometric data; however, their scattered distribution and shallow depth suggest they are likely associated with small metallic objects in the near-surface layer - such as the previously identified wires, fragments of cookware, or bottles. The only anomaly for which a clear and consistent correlation between resistivity and magnetometric data exists is the aforementioned mound, whose presence is likely linked to anthropogenic surface modification.

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*Fig.5.3. Distribution of subsurface electrical resistivity at depths of 0.25, 0.5, 0.8, 1.1, 1.6, and 2.3 m within the boundaries of survey area no. 5.*

* Fig.5.4. Distribution of subsurface electrical resistivity at depths of 0.25, 0.5, 0.8, 1.1, 1.6, and 2.3 m within the boundaries of survey area no. 5, with the magnetic anomalies marked in dashed red lines.*

1. **Summary**

The geophysical surveys conducted using gradient magnetometry and electromagnetic conductivity methods enabled the identification of several structures of potential archaeological significance, as well as the detection of features and phenomena associated with signal interference. The character and intensity of the recorded anomalies varied significantly between the analyzed polygons, which can be attributed both to local geological and geomorphological conditions and to the presence of interfering elements such as metal fences or scattered surface debris.

The most distinct and spatially coherent anomalies were recorded in Polygons 1 and 4, where data from both methods suggest the presence of structures with a potentially anthropogenic origin. In particular, for Polygon 1, the correlation between magnetic and resistivity results allows for the preliminary identification of an anomaly that may correspond to a mass grave. Similarly, in Polygon 4—despite the weaker magnetic effects—the conductivity data revealed the presence of features that may be relevant from an archaeological prospection perspective.

Polygons 2 and 3 were affected by interference related to the proximity of fencing and varying degrees of data noise; however, localized correlations between the two methods were still observed. In the case of Polygon 5, due to extensive surface litter and an unusual topographic configuration, data interpretation proved most challenging. Only within the area of an earthen mound in the northeastern part of the polygon was a consistent conductivity anomaly recorded, partially confirmed by magnetometric data.

The report was prepared on March 27, 2024.

dr inż. Szymon Oryński